

The Santorini Inflation Episode, Monitored by InSAR and GPS

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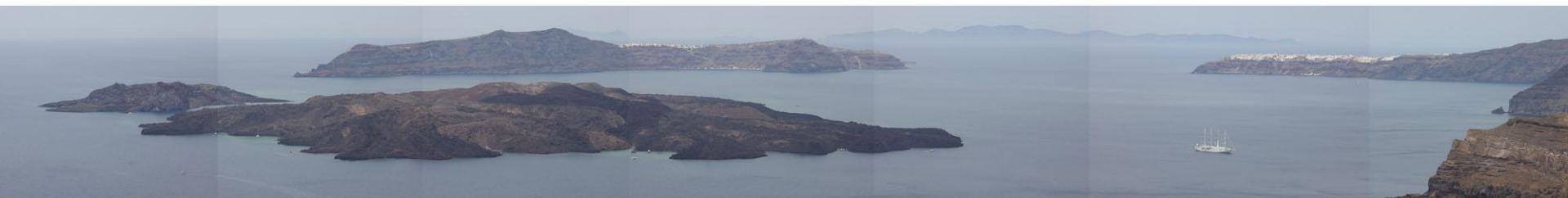
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⁴ The University of Alabama, Department of Geological Sciences, USA

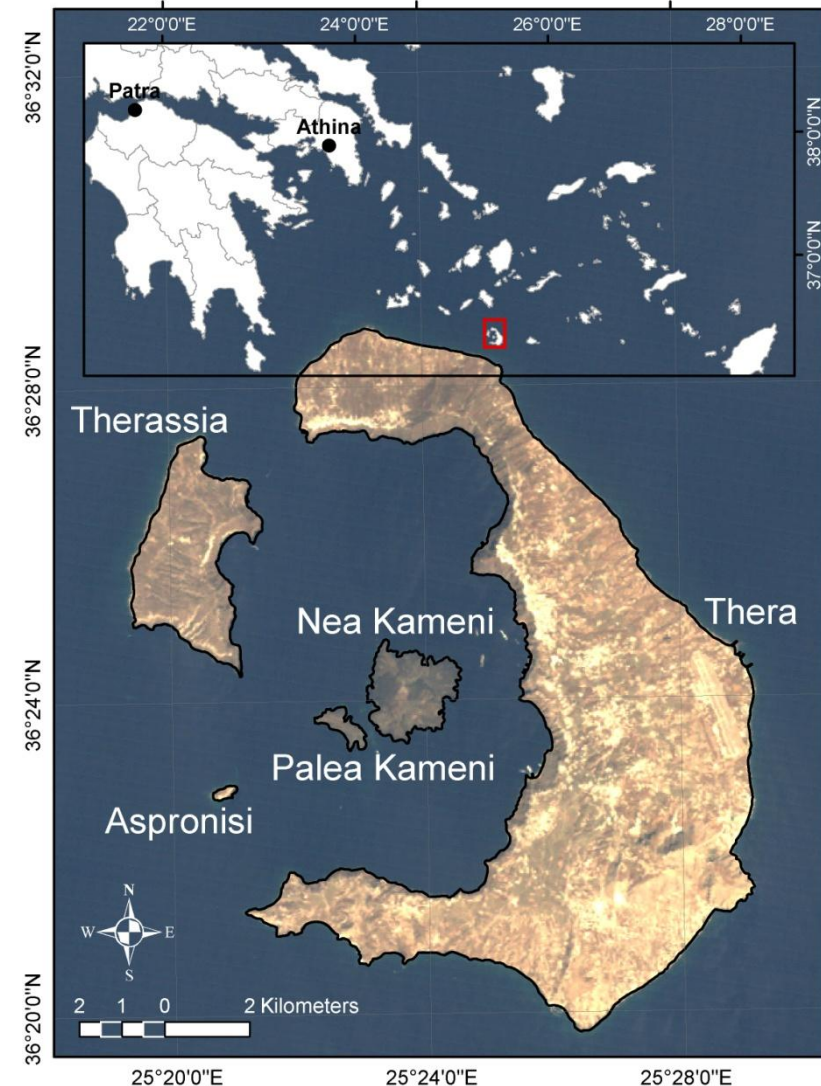
⁵ Institute of Geodynamics / National Observatory of Athens, Greece



Background information on Santorini



- Santorini Volcanic Complex is the most active part of the South Aegean (Hellenic) Volcanic Arc.
- Several eruptions led to the present form of the Kameni islands (197 BC, 46 AD, 726, 1570, 1707, 1866, 1925, 1939, 1950)
- Most recent seismic sequence ended in 1950
- Since then, Santorini volcano has been in a 'quiet' phase, with insignificant deformation (confirmed by GPS and InSAR)





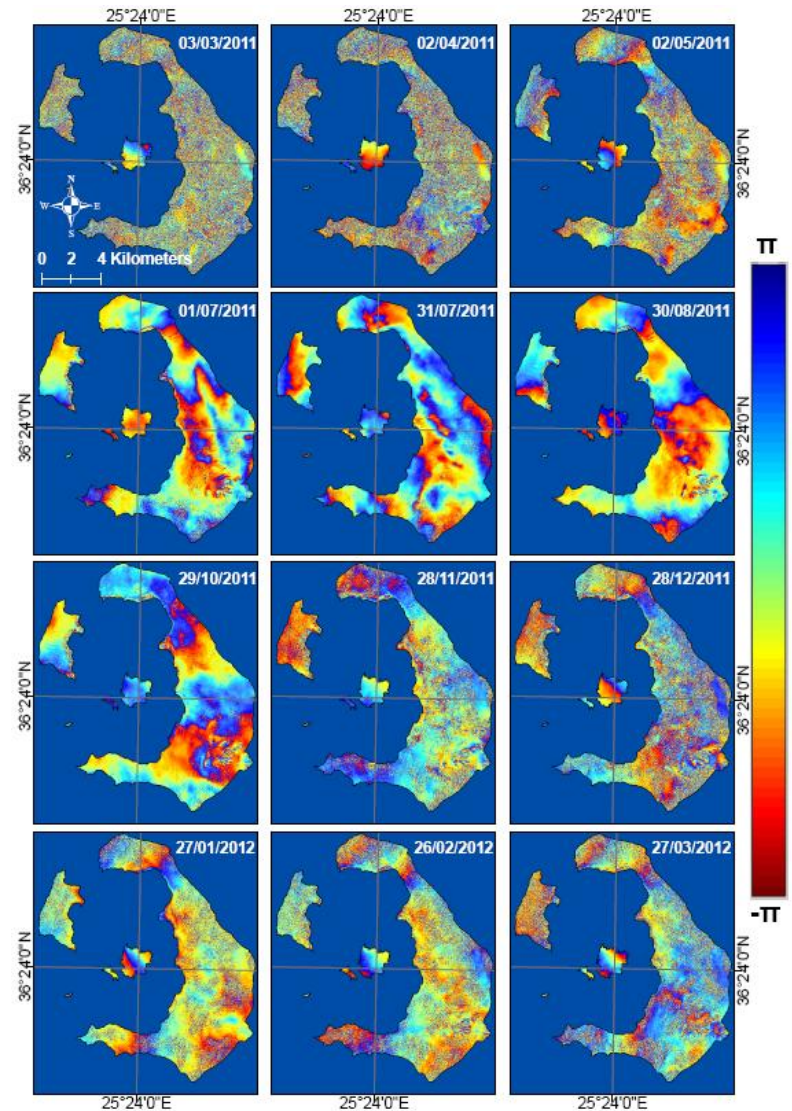
- Newman et al., Geophysical Research Letters, March 2012
 - Conducted GPS campaigns to quantify unrest for the first time
 - Modeled volcanic source using a Mogi model
- Parks et al., Nature Geoscience, Sept. 2012
 - Used stacked Envisat and TerraSAR-X interferograms
 - Concluded that shallow magma chamber is charged episodically by high-flux batches of magma in Santorini
- Papoutsis et al., Geophysical Research Letters, Jan. 2013
 - Applied PSI and SBAS on ENVISAT data
 - Analyzed data from 10 cGPS stations
 - Claimed that the unrest episode has ended
- Subsequent pubs (Parks et al./EPSL/July2013; Feuillet/GRL/July2013; Foumelis et al./GJI/April2013; Lagios et al./Tectonophysics/March2013; Tassi et al./Bul. Of Volcanology /March2013, Chouliaras et al./NHES/April/2012)

Input data and methodology

Satellite interferometry – PSI & SBAS



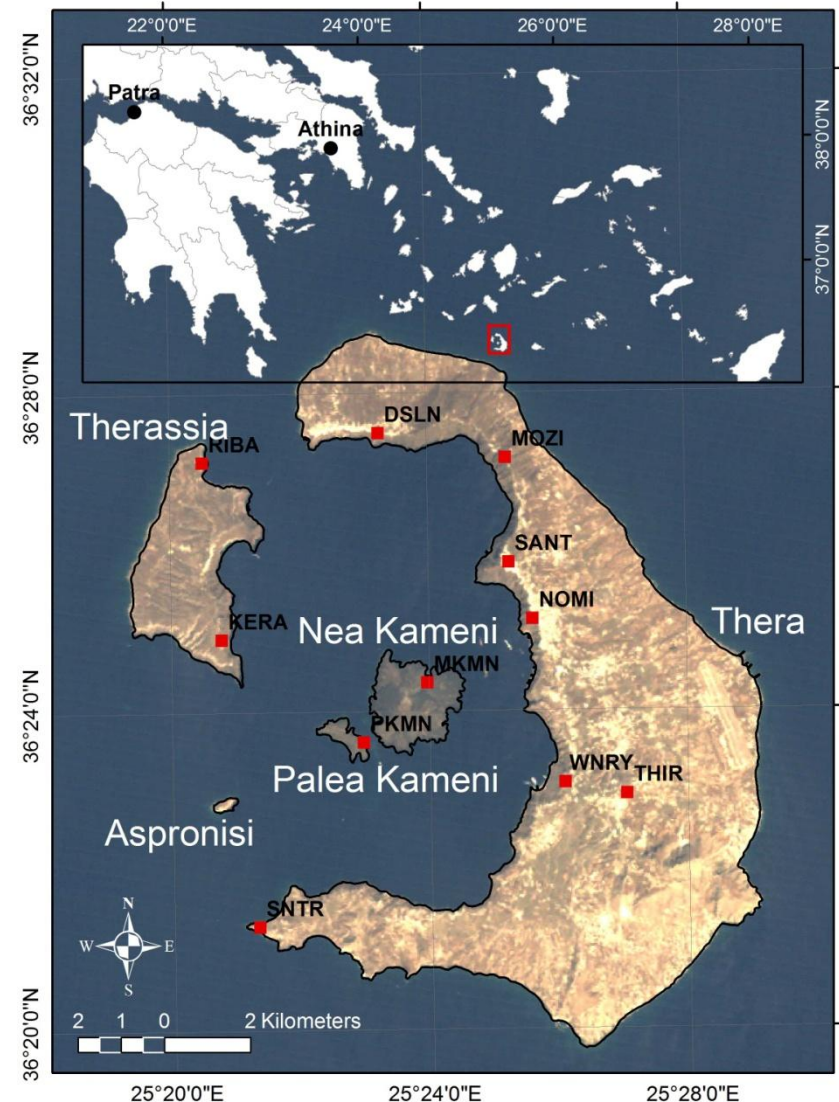
- 13 ASAR ENVISAT, descending mode
- Last orbit before the end of the mission in April 2012
- Short spatial & temporal baselines
- Swath I6, leading to increased sensitivity to the E-W horizontal components
- S/W: Gamma, ROI_PAC, DORIS, StaMPS (Hooper et al., JGR, 2007)
- PSI challenging due to the limited number of scenes
- Oversampling by a factor of 2





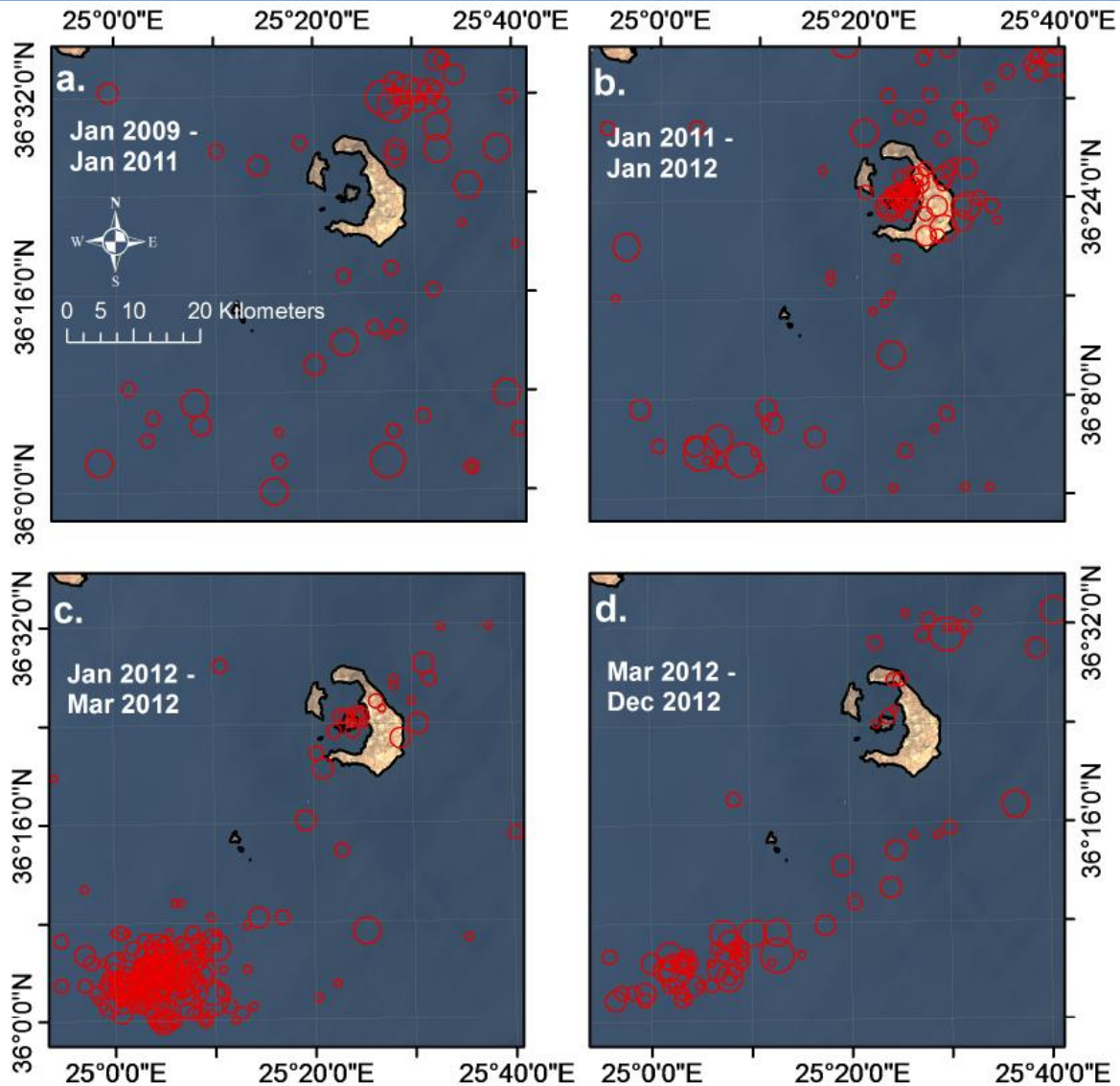
Receivers installed and maintained by:

- UNAVCO
- NTUA
- Georgia Tech/University of Patras
- COMET/University of Oxford
- NOANET/NKUA



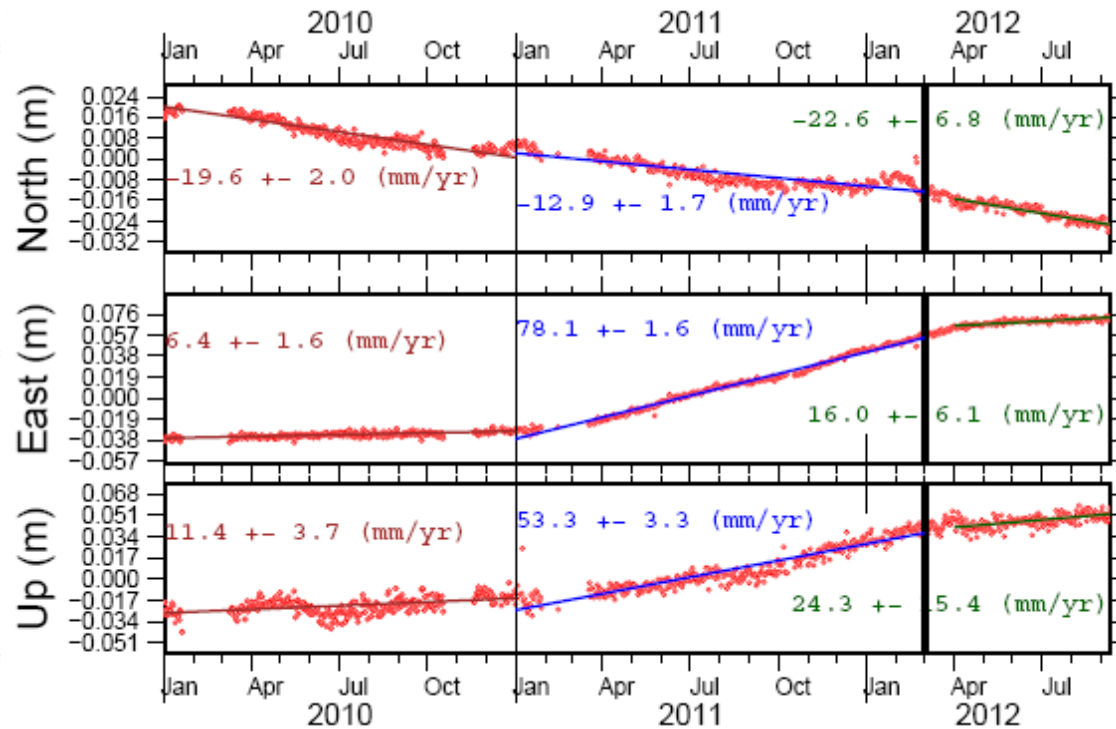
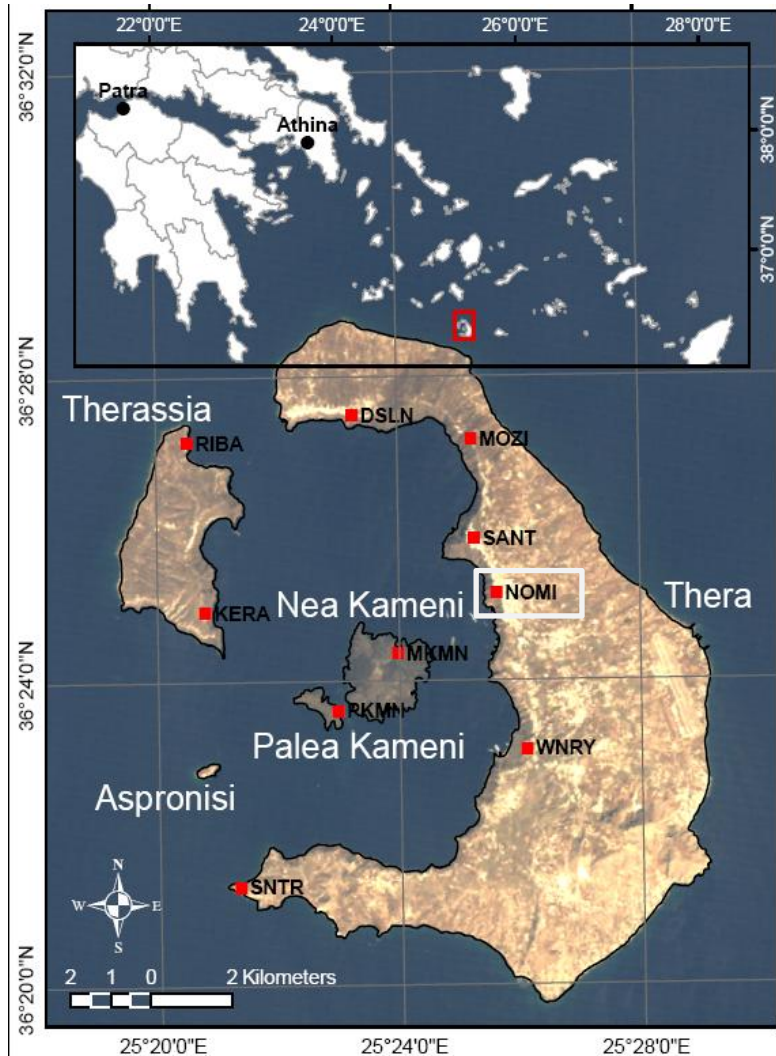
The 2011-2012 unrest

Seismicity



The 2011-2012 unrest

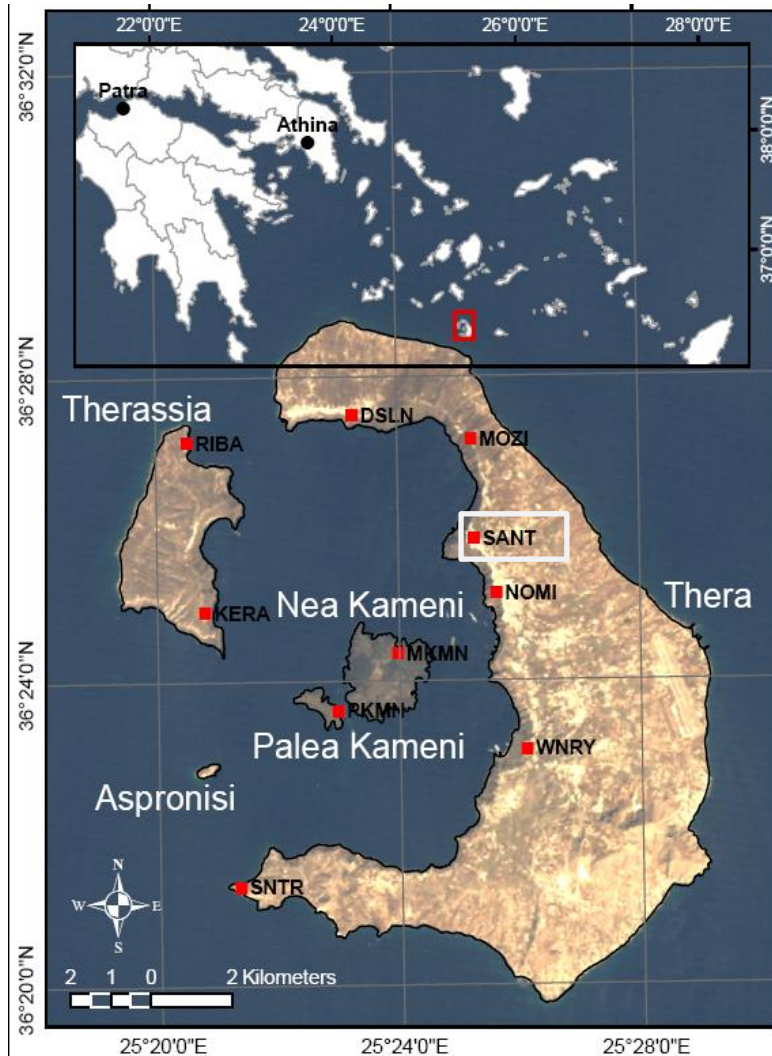
Deformation field - GPS



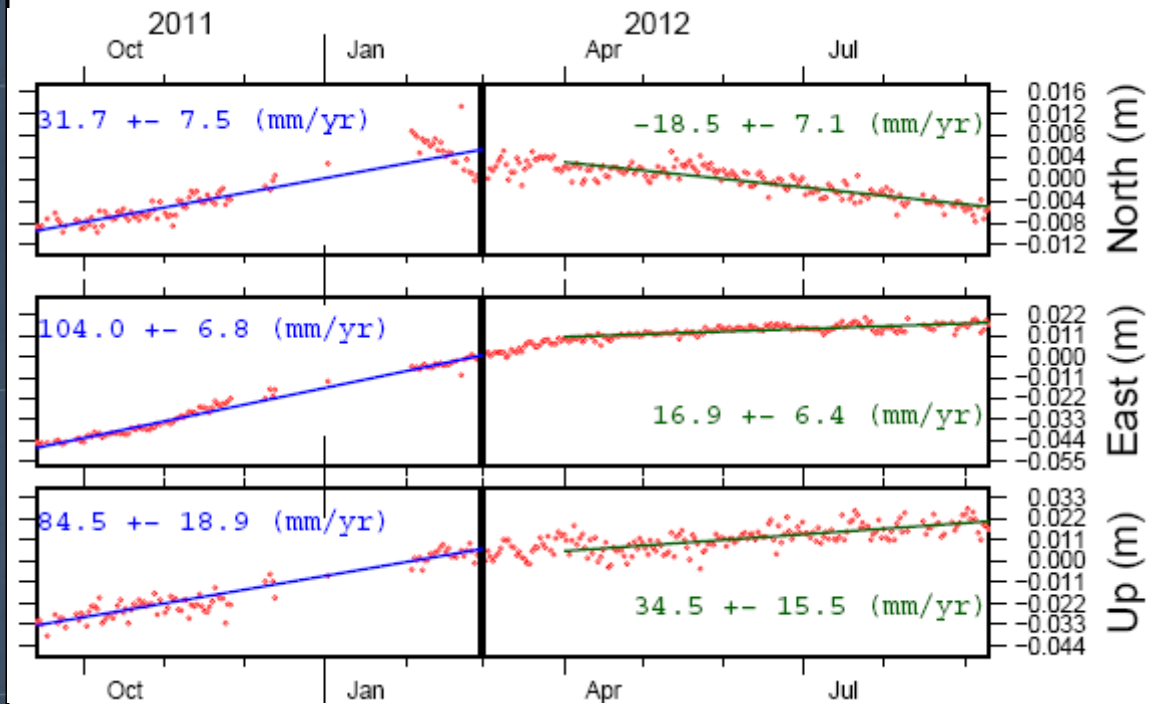
NMI station

The 2011-2012 unrest

Deformation field - GPS



SANT station

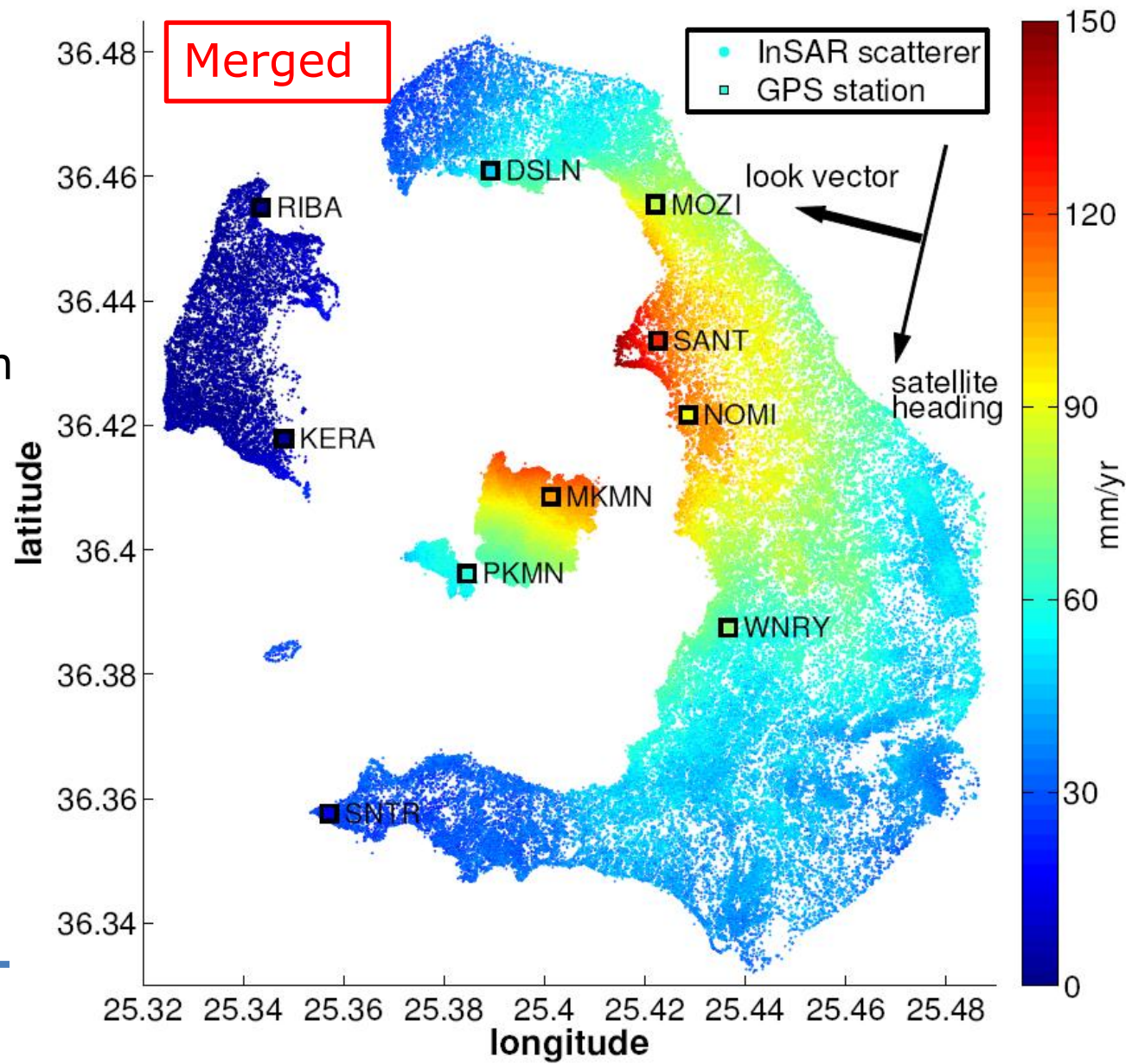


The 2011-2012 unrest

Deformation field – PSI & SBAS



- Merged rates from PSI and SBAS (Hooper, GRL, 2008)
- Identified more than 250000 coherent pixels
- Radially decaying deformation pattern
- 150 mm/yr maximum displacement rate



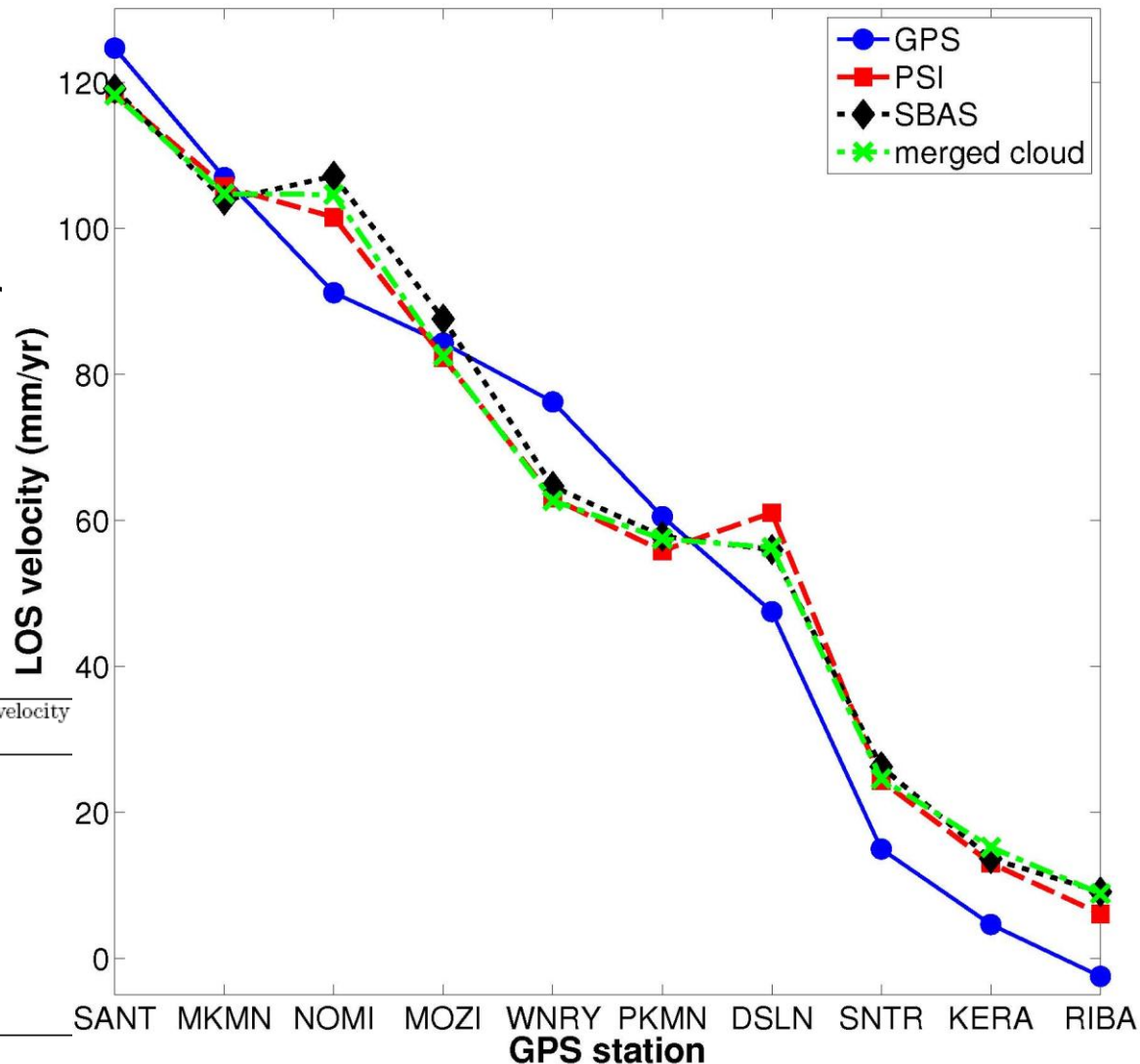


RMS differences

GPS-PSI: 8.72mm/yr

GPS-SBAS: 9.28 mm/yr

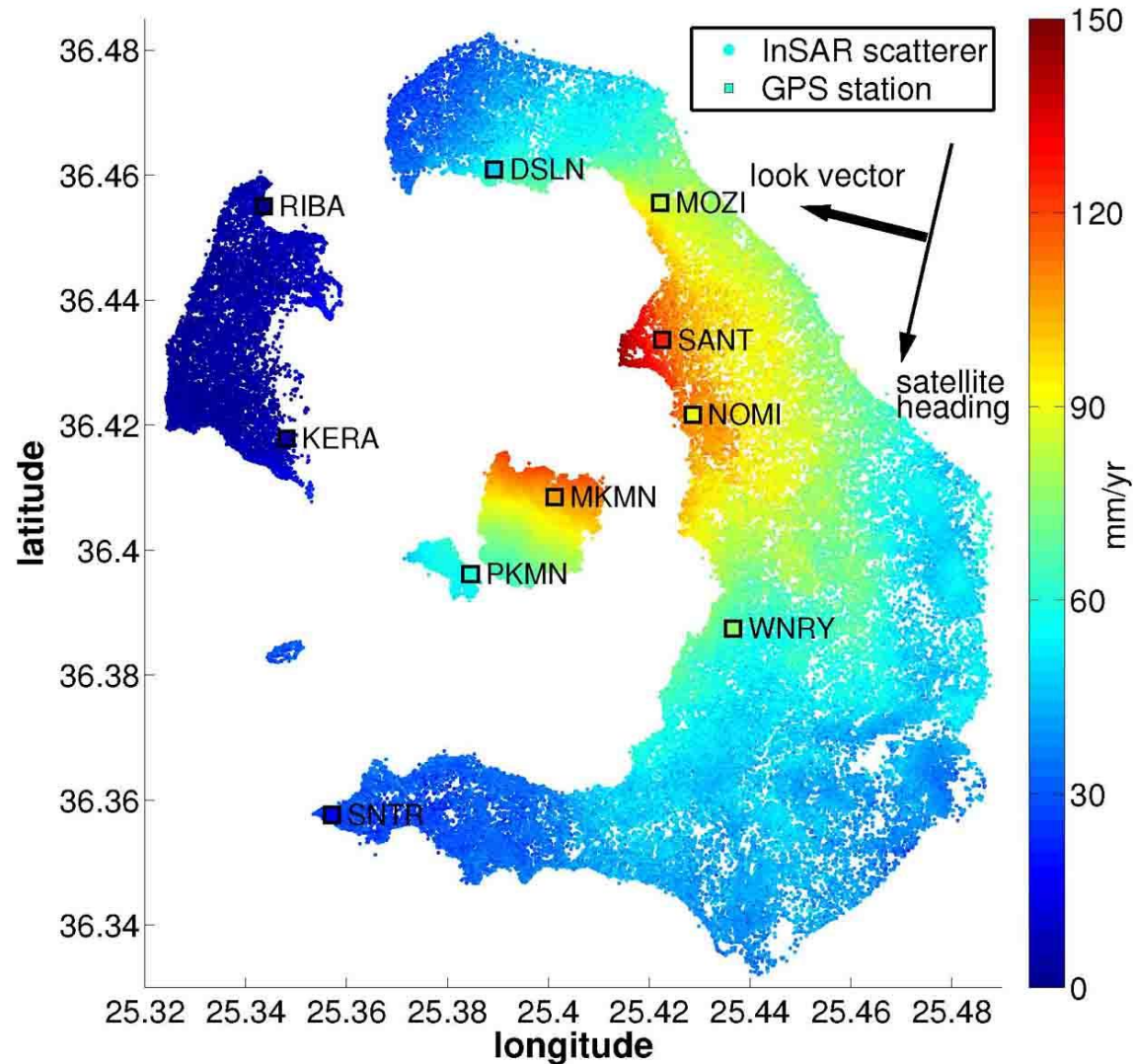
GPS-MERGED: 9.12 mm/yr



* merged PSI and SBAS cloud



InSAR analysis



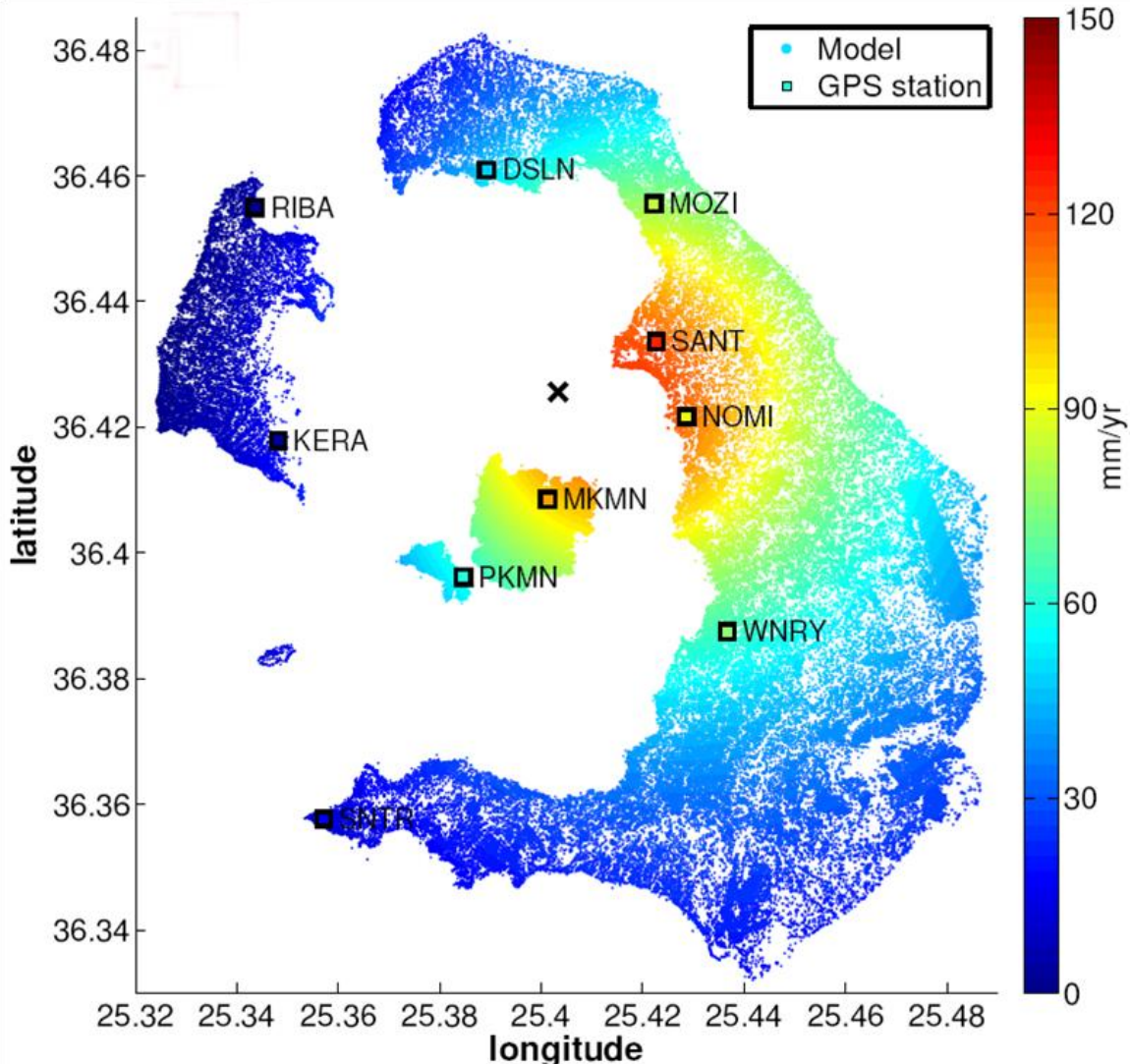
The 2011-2012 unrest

Mogi model - InSAR



x marks the location of the Mogi source inferred from InSAR

Mogi model



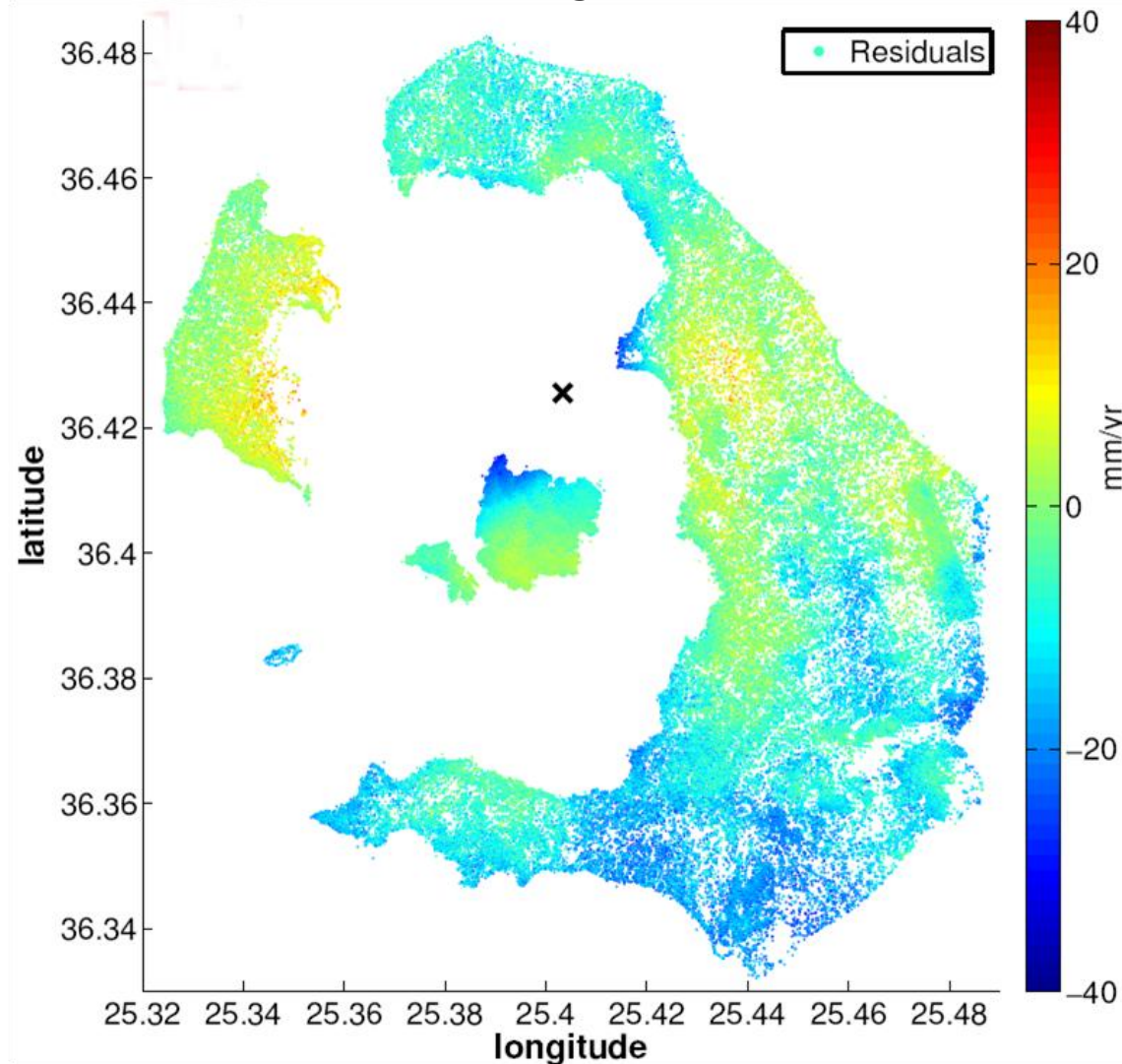
The 2011-2012 unrest

Mogi model - InSAR



x marks the location of the Mogi source inferred from InSAR

Residuals
(model-InSAR)

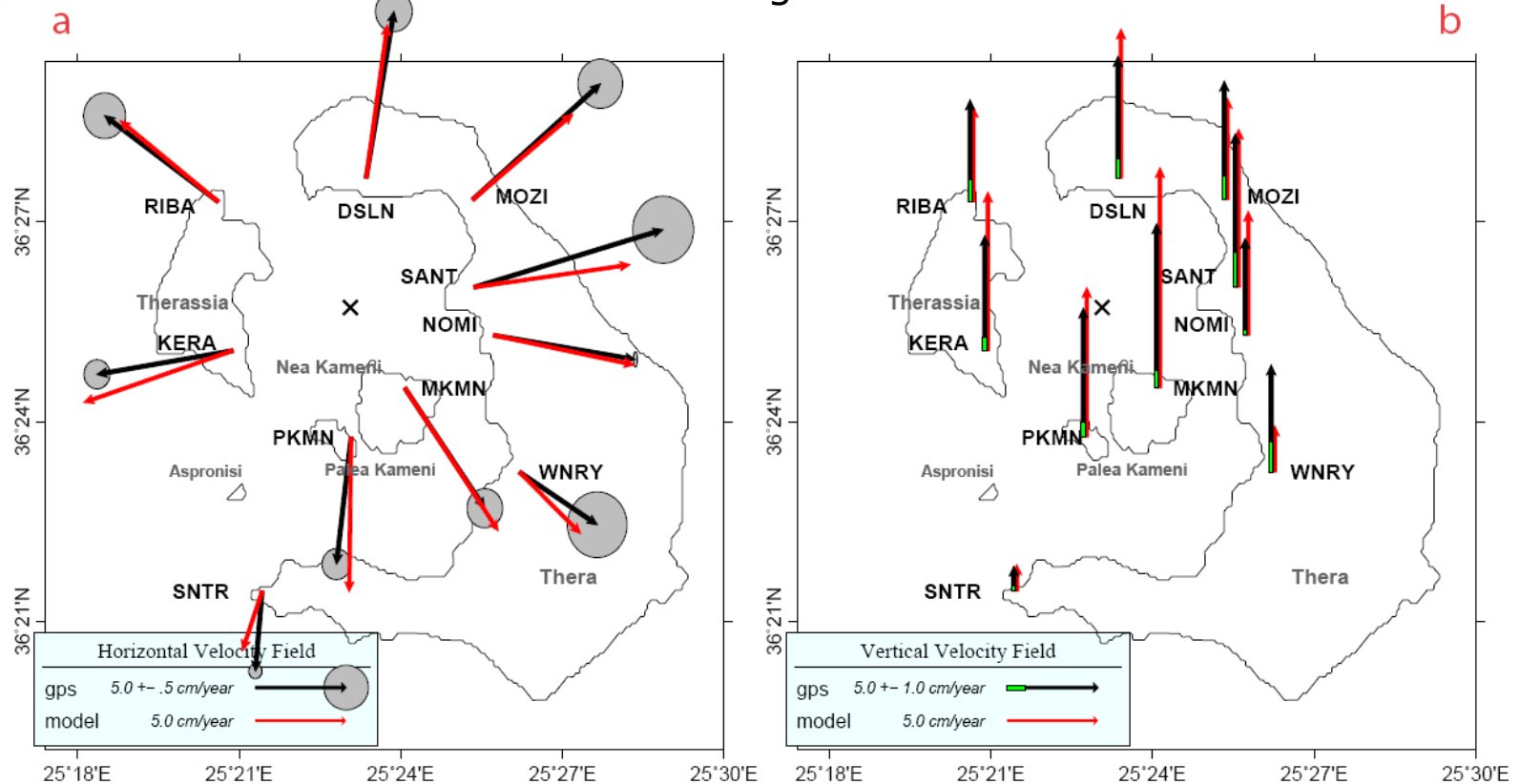


The 2011-2012 unrest

Mogi model - GPS



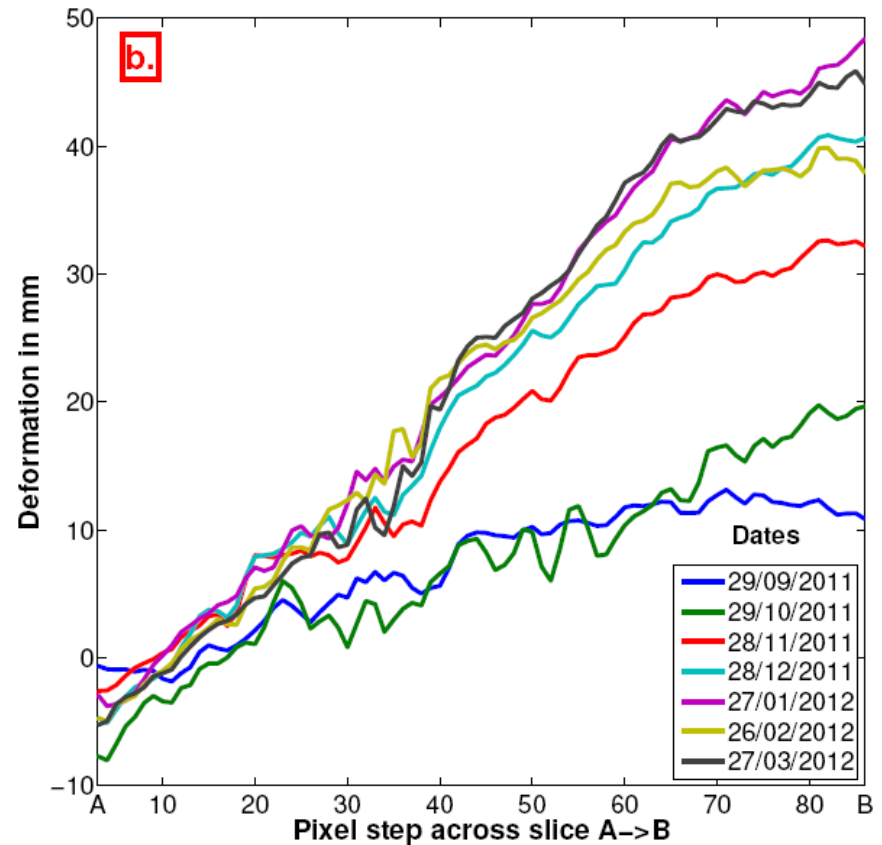
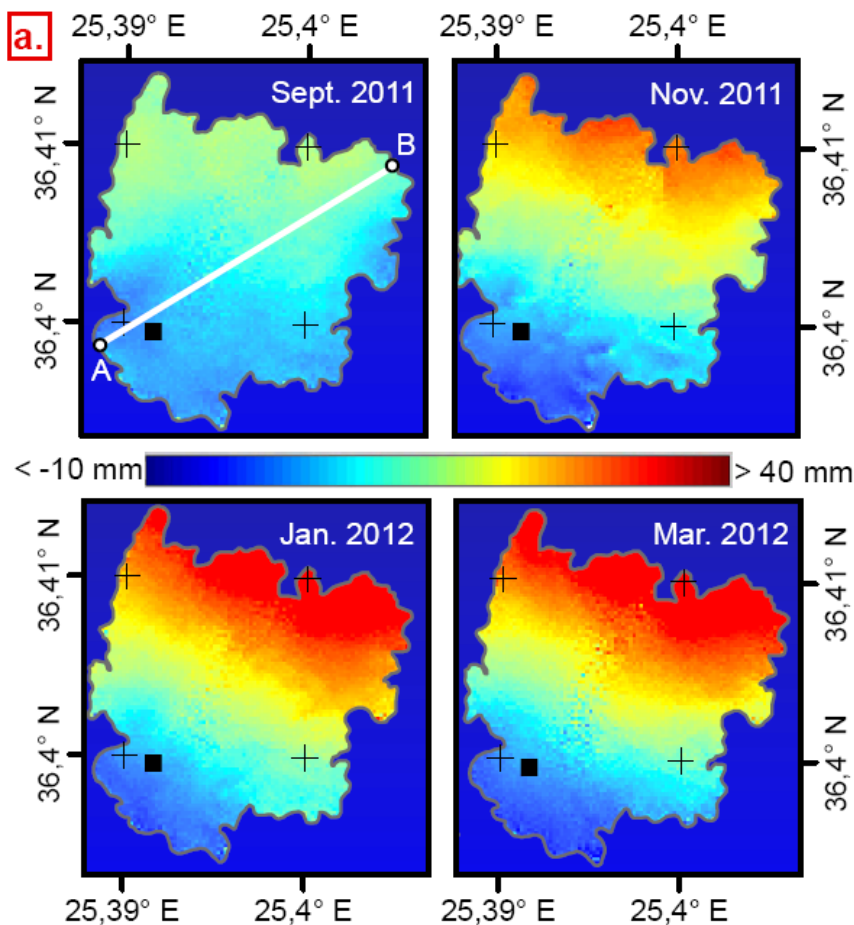
x marks the location of the Mogi source inferred from GPS



Mogi model best fit parameters for the GPS and InSAR data

Data set	Longitude	Latitude	Depth/km	$\Delta V/10^6 \text{ m}^3/\text{yr}$	χ^2/dof^a
3-component GPS	25.3844	36.4286	$3.48^{+0.19}_{-0.17}$	$12.4^{+0.9}_{-0.8}$	9.1
InSAR	25.4033	36.4256	$6.28^{+0.02}_{-0.02}$	$24.2^{+0.1}_{-0.1}$	3.52

^aDegrees of freedom.



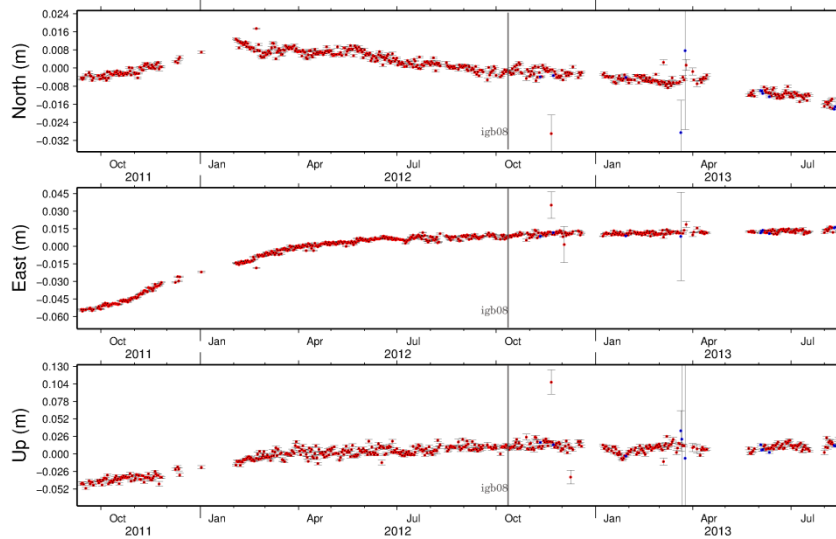
(a.) Unwrapped differential interferograms zoomed in the Nea Kameni region with reference to 03/2011. While the magnitude of uplift clearly increases for the first three interferograms, in 03/2012 the deformation is similar to the one observed in 01/2012. (b.) Cumulative deformation in millimeter across slice AB for selected Envisat acquisition dates.

The end of the episode

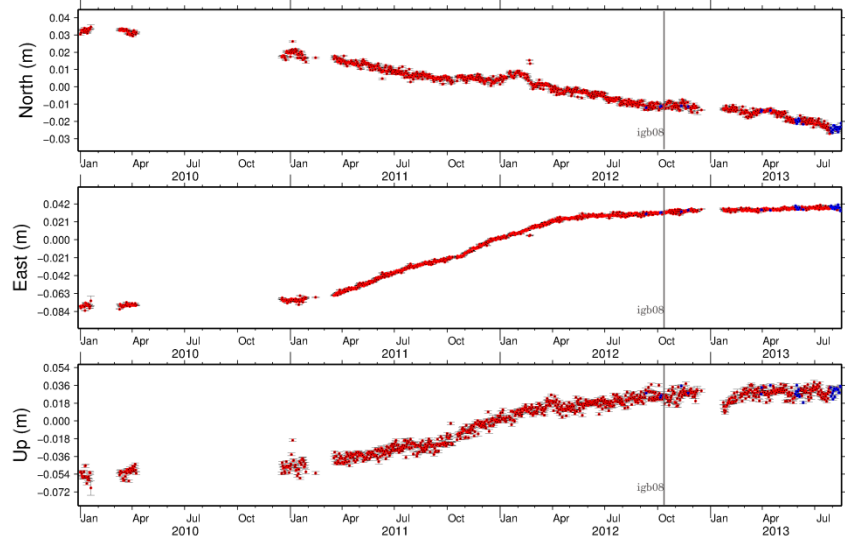
GPS – Raw data



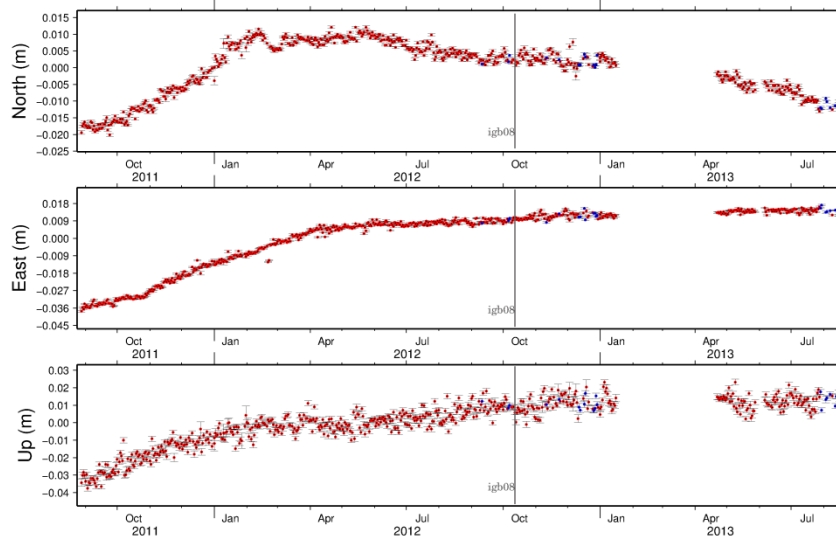
cGPS station sant



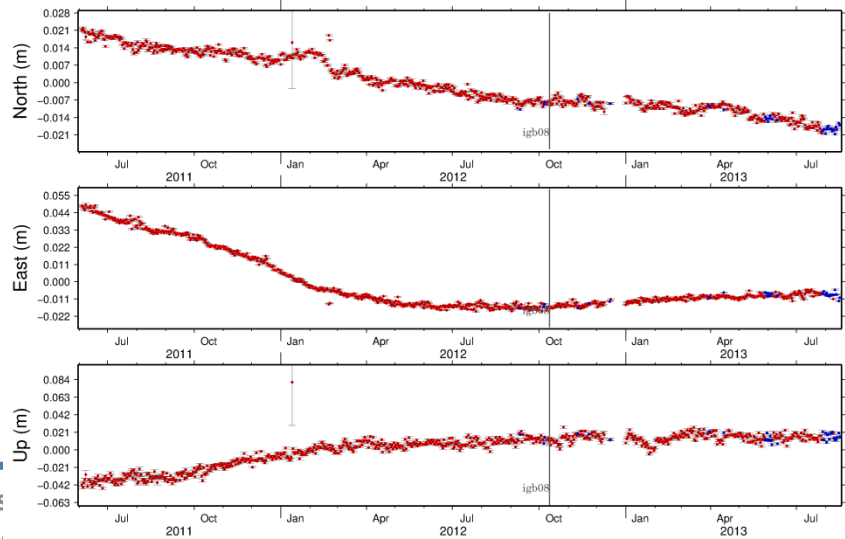
cGPS station nomi



cGPS station mozi



cGPS station kera



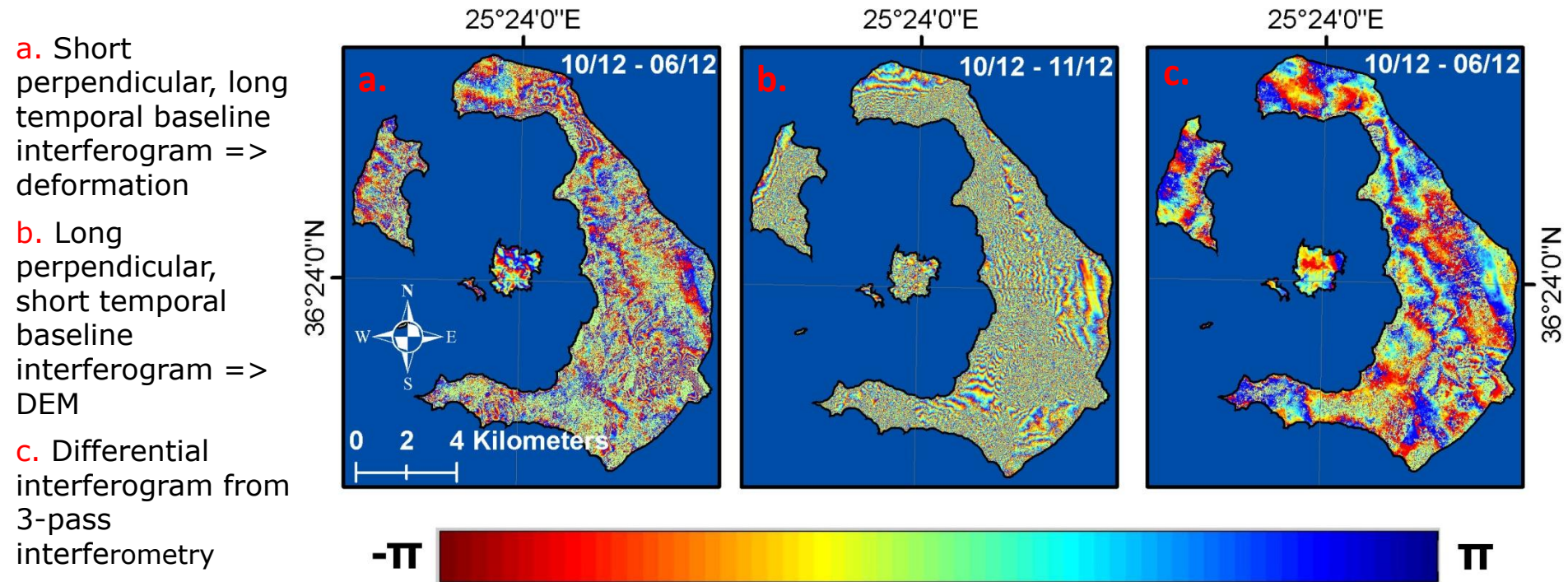


- Mogi source model seems to be suitable (in agreement with Newman et al., Parks et al.)
- Unless a very deep hydrothermal fluid reservoir exists beneath the caldera, this episode was likely to be one of magmatic inflation of the shallow chamber
- Inflation has diminished since the end of February 2012
 - New phase of relative stability
 - Reduced probability for an imminent volcanic eruption

- Daily GPS solutions for Santorini

➤ http://dionysos.survey.ntua.gr/src/cgps_processing_main.htm

Ongoing work with COSMO-SkyMed SAR data (3-pass interferometry)





- BEYOND project aims at establishing a Centre of Excellence for Earth Observation based monitoring of Natural Disasters in south-east Europe



- <http://www.beyond-eocenter.eu/>
- June 2013 – 2016, €2.3M EU contribution
- Beneficiary is the National Observatory of Athens
- In the framework of BEYOND we will:
 - Set up innovative integrated observational solutions to allow a multitude of monitoring networks (space borne and ground-based) to operate in a complementary, unified and coordinated manner
 - Create archives and databases of long series of observations and derived higher level products
 - Collaborate with key players in Europe for geophysical research
 - Recruit experienced researchers and upscale existing s/w and h/w capacities



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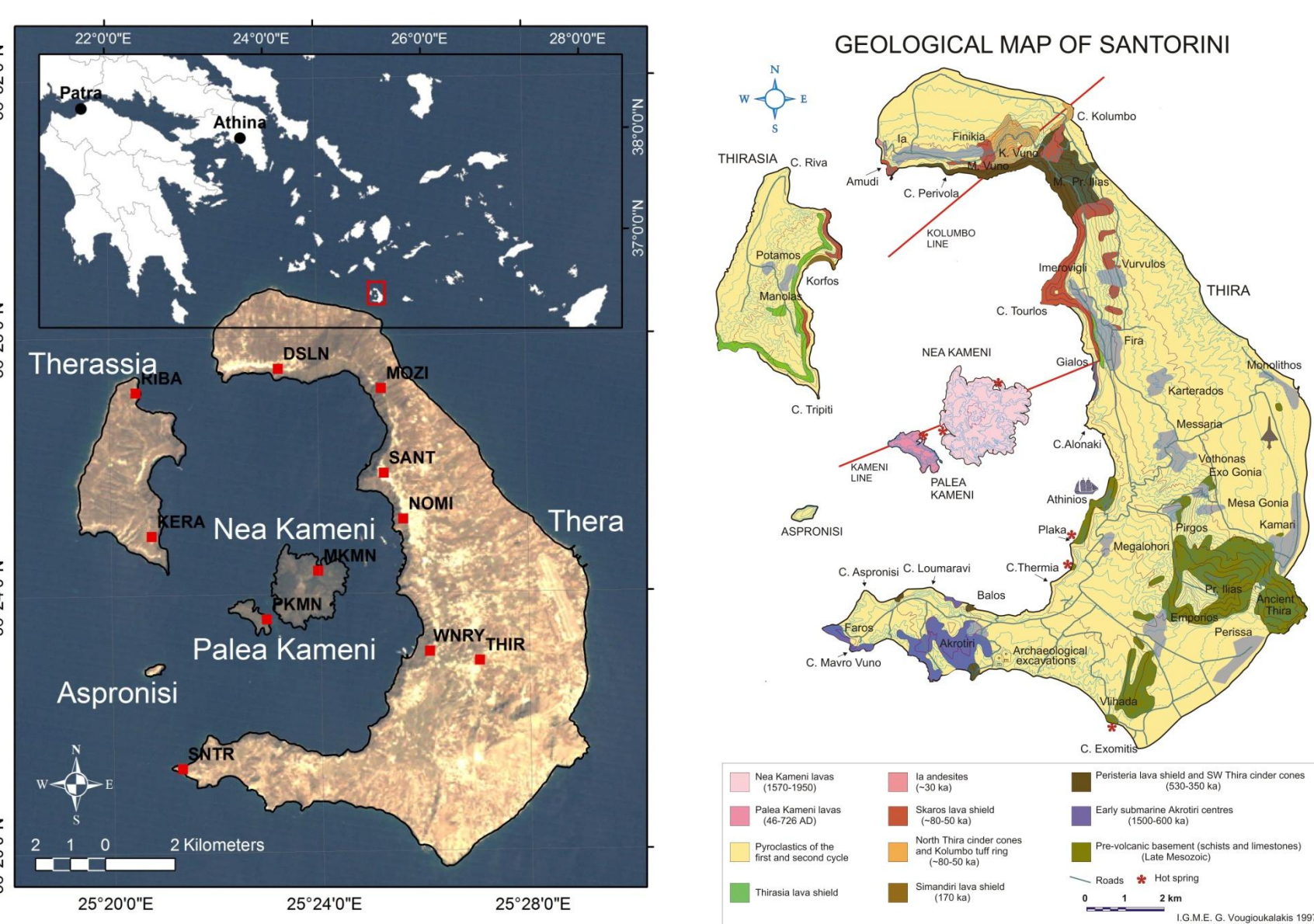


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² Dionysos Satellite Observatory / National Technical University of Athens, ³ Department of Earth, Atmospheric and Planetary Sciences, Massachusetts Institute of Technology, ⁴ The University of Alabama, Department of Geological Sciences, USA, ⁵ Institute of Geodynamics / National Observatory of Athens

Background

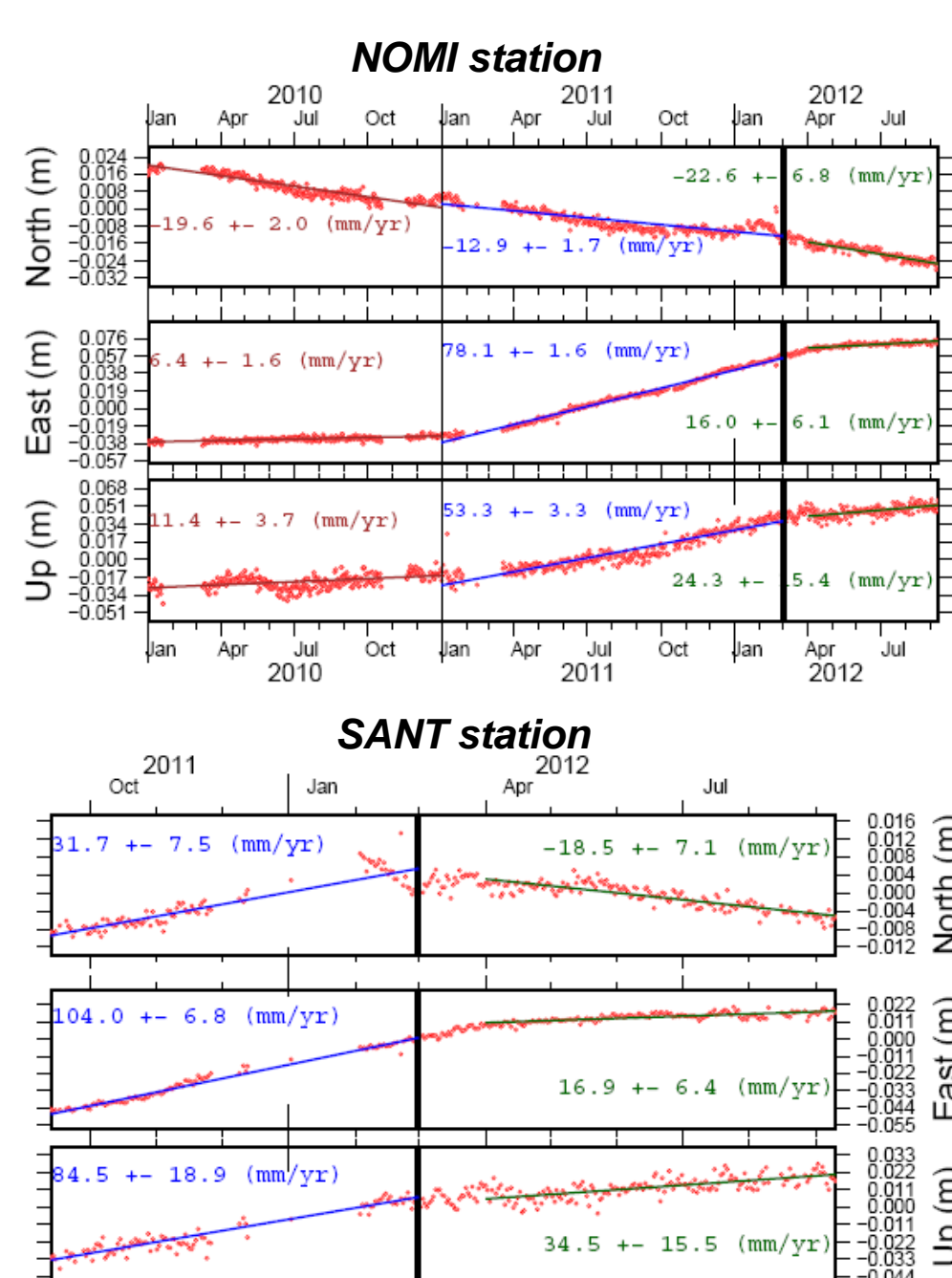
- ✓ The Santorini volcanic complex is comprised of four islands: Therassia and Thera islands form the caldera rim; Palea Kameni and Nea Kameni have built up in the central caldera
- ✓ Santorini Volcanic Complex is the most active part of the South Aegean (Hellenic) Volcanic Arc.
- ✓ Several eruptions led to the present form of the Kameni islands
- ✓ Most recent seismic sequence ended in 1950
- ✓ Since then, Santorini volcano has been in a 'quite' phase,



(left) Landsat 5 TM scene depicting Santorini and the locations of the installed cGPS (right) Simplified geological map of Santorini (Vougioukalakis, 1997)

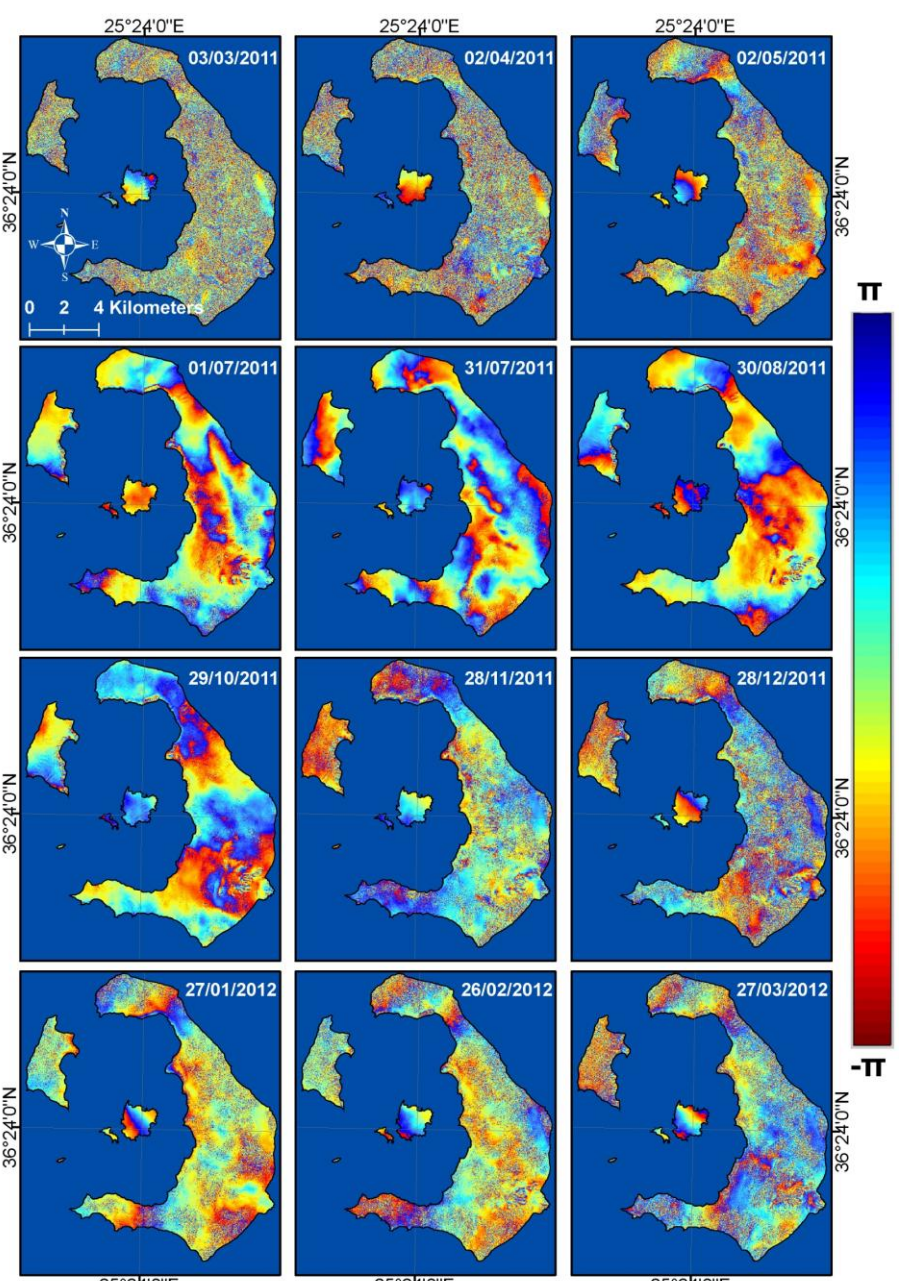
Input data and methodology

GPS



Raw GPS time series for selected stations

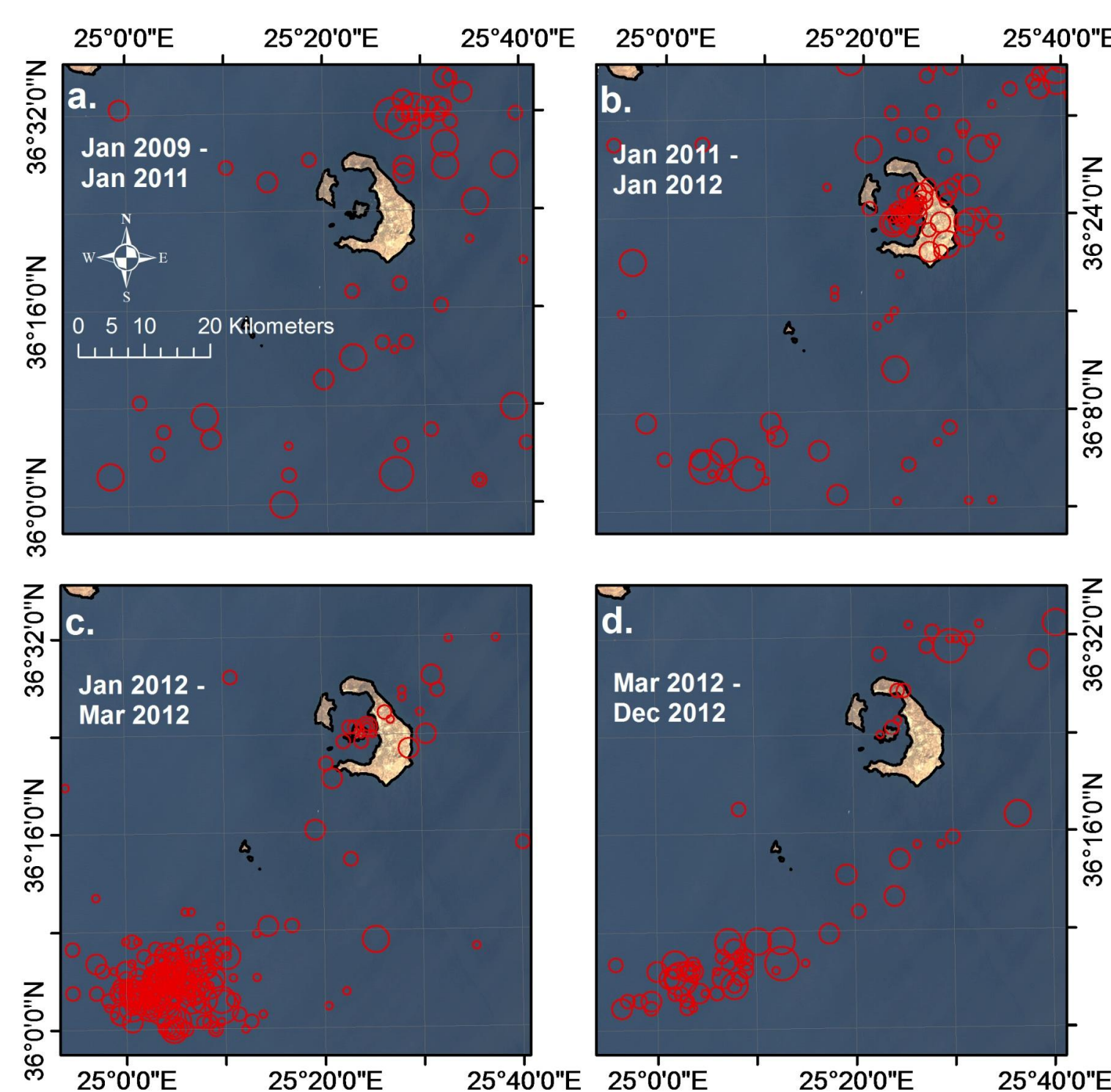
DInSAR



Wrapped Envisat interferograms with respect to the September 2011 acquisition

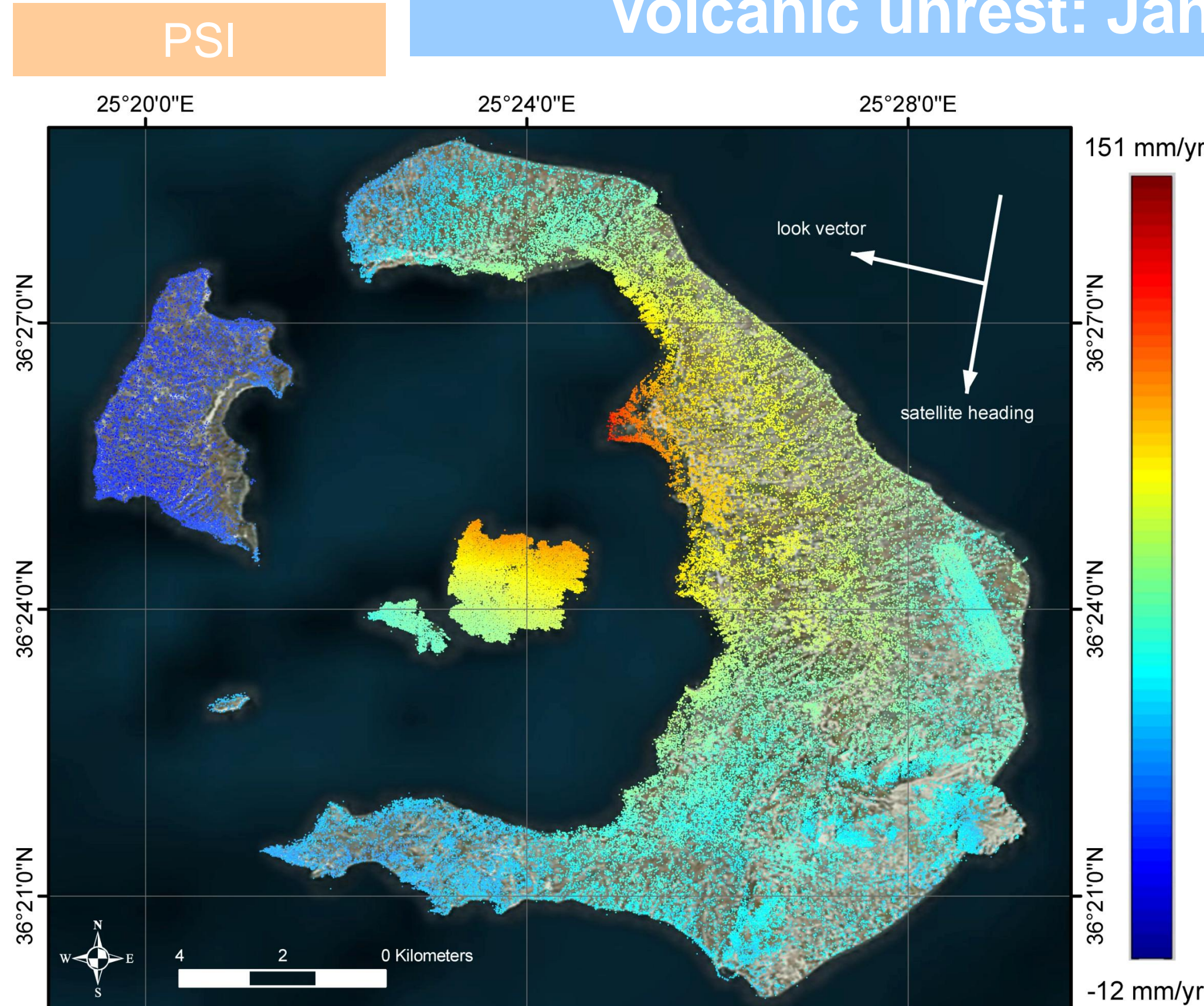
- ✓ 13 ASAR Envisat descending mode
- ✓ Last orbit before the end of the mission in April 2012
- ✓ Short spatial & temporal baselines
- ✓ Swath I6, leading to increased sensitivity to the E-W horizontal components
- ✓ S/W: Gamma, StaMPS (Hooper et al., JGR, 2007)
- ✓ PSI challenging due to the limited number of scenes

Seismicity

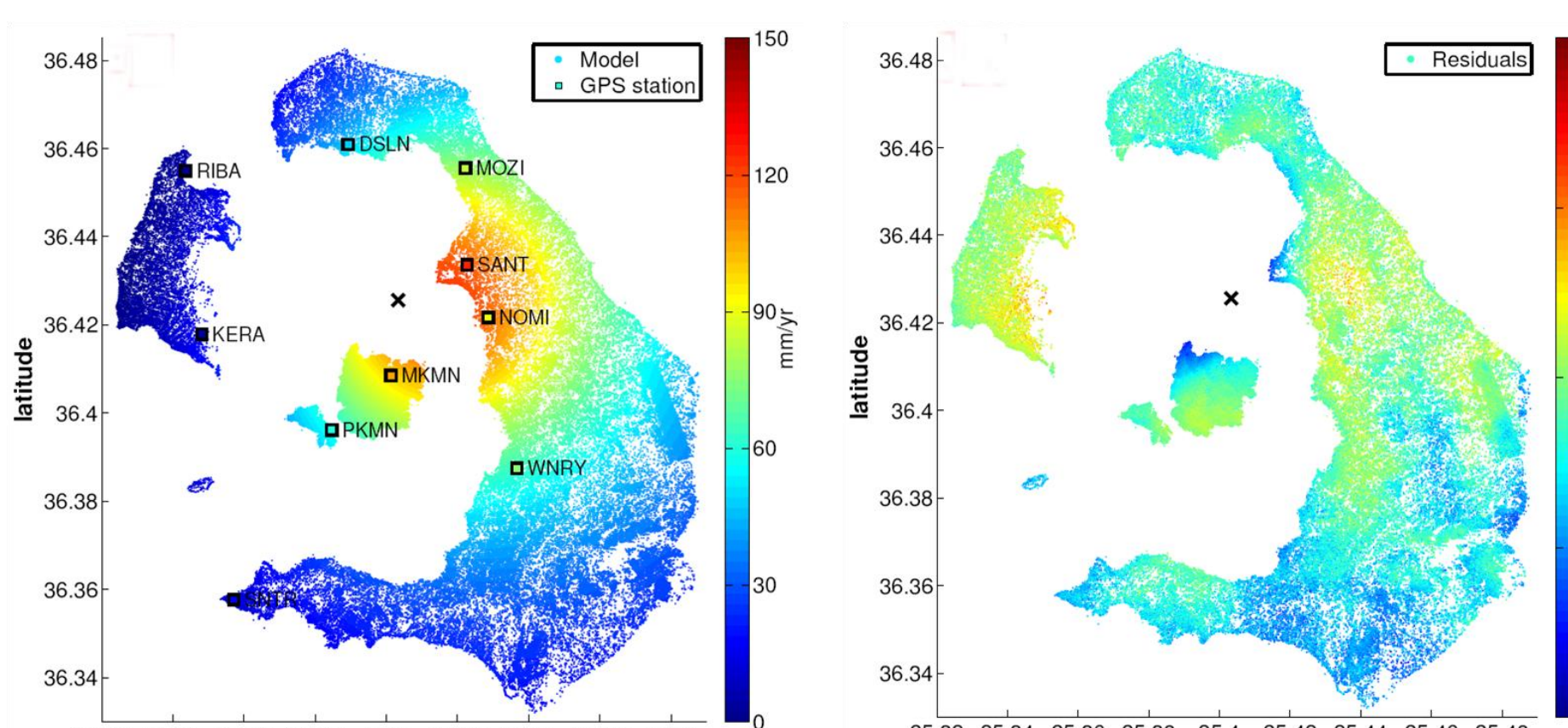


Regional seismicity for selected time spans

Volcanic unrest: January 2011



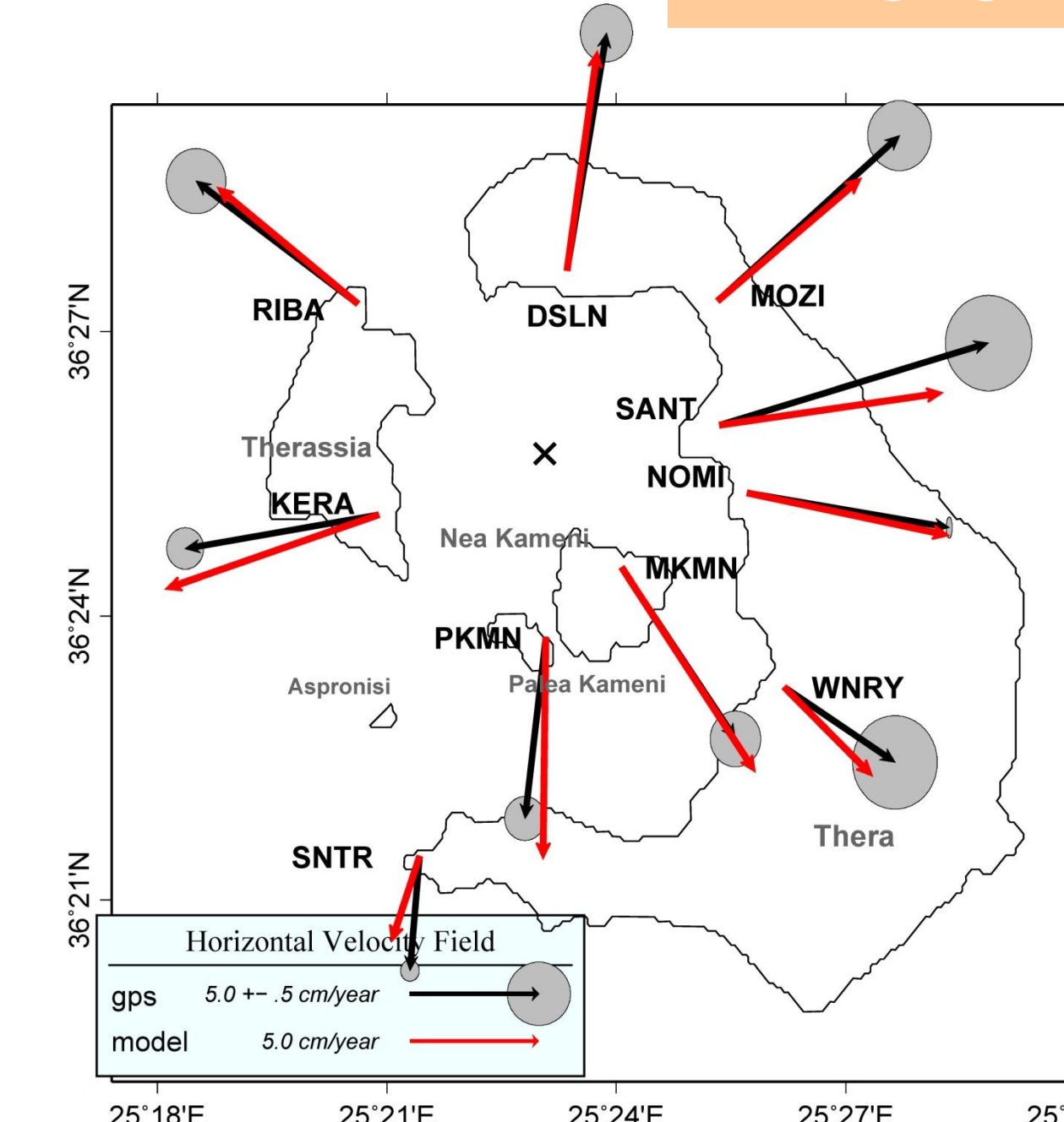
Observed LOS deformation rates from Jan. 2011- Feb. 2012



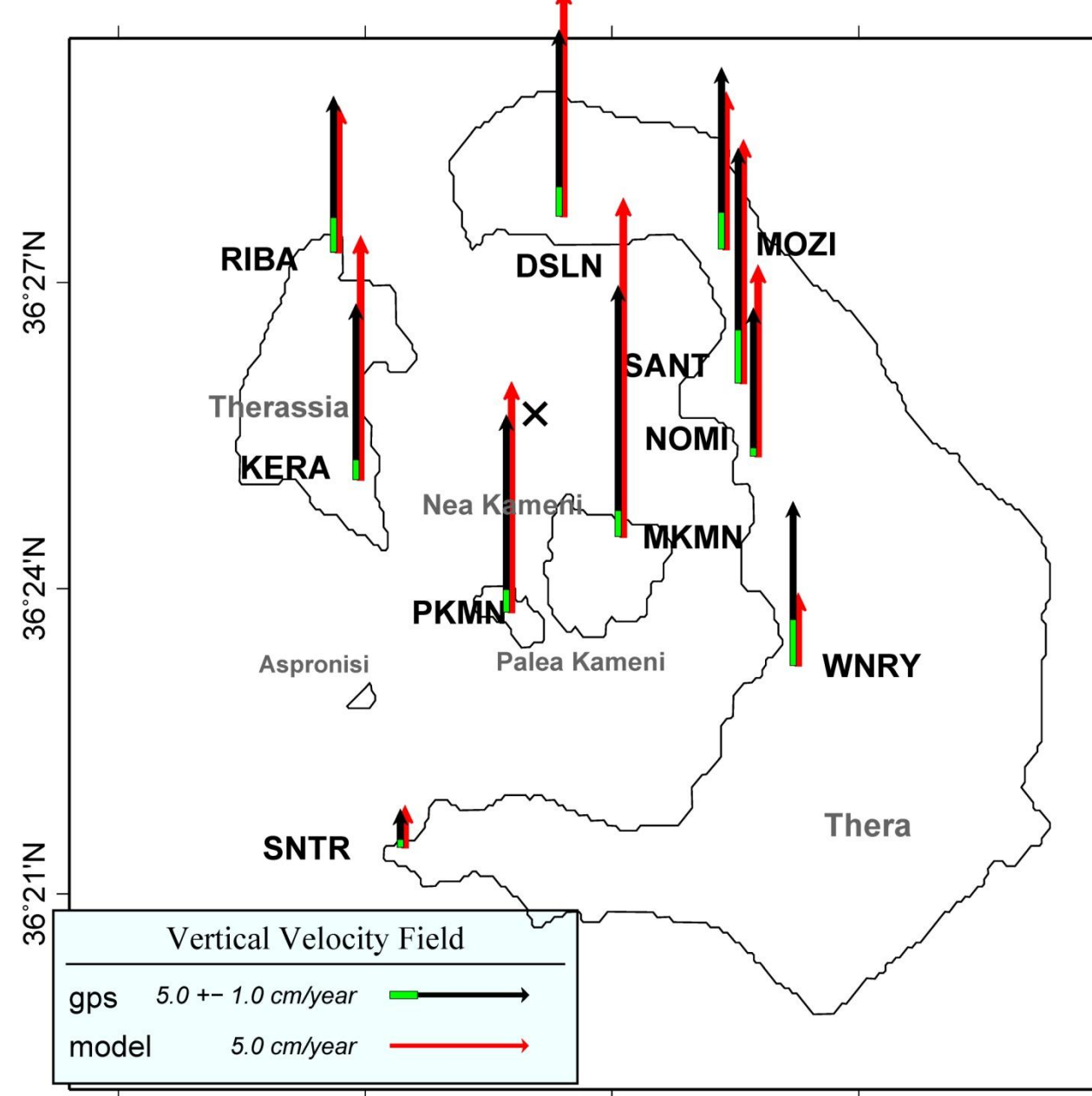
Deformation from best fit Mogi source (location marked with an x)

Residuals: PSI minus Mogi model

GPS



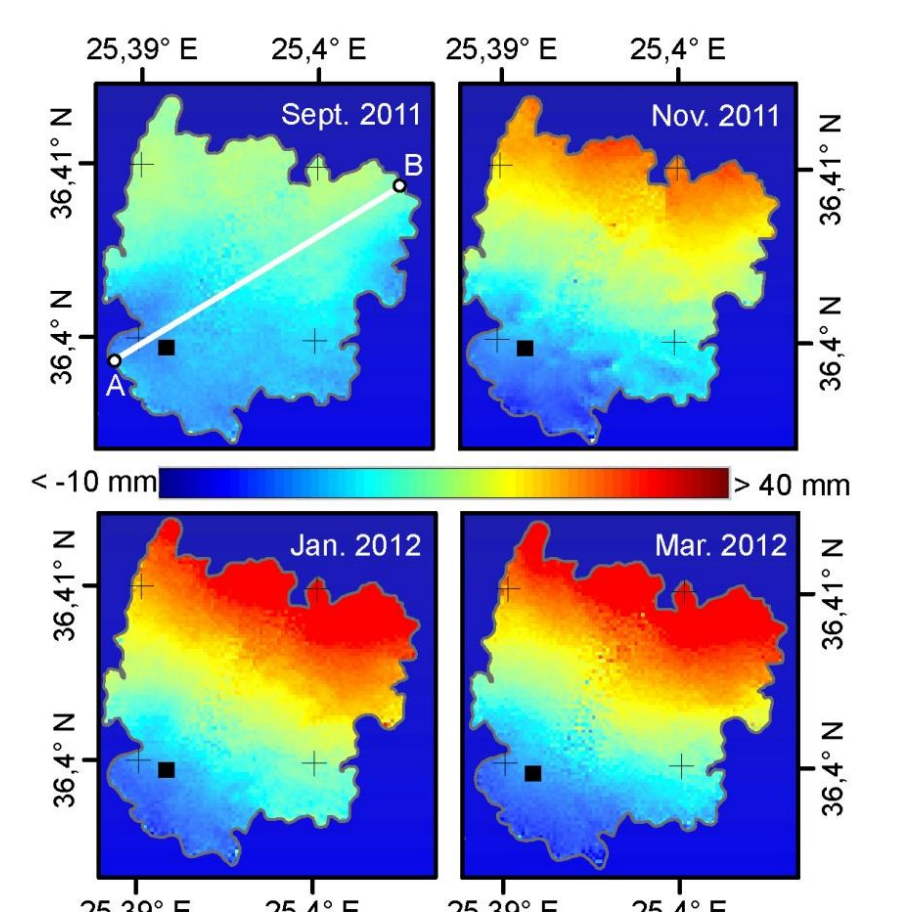
Horizontal GPS velocities and velocities from a best fit Mogi source (location marked with an x)



Vertical GPS velocities and velocities from a best fit Mogi source

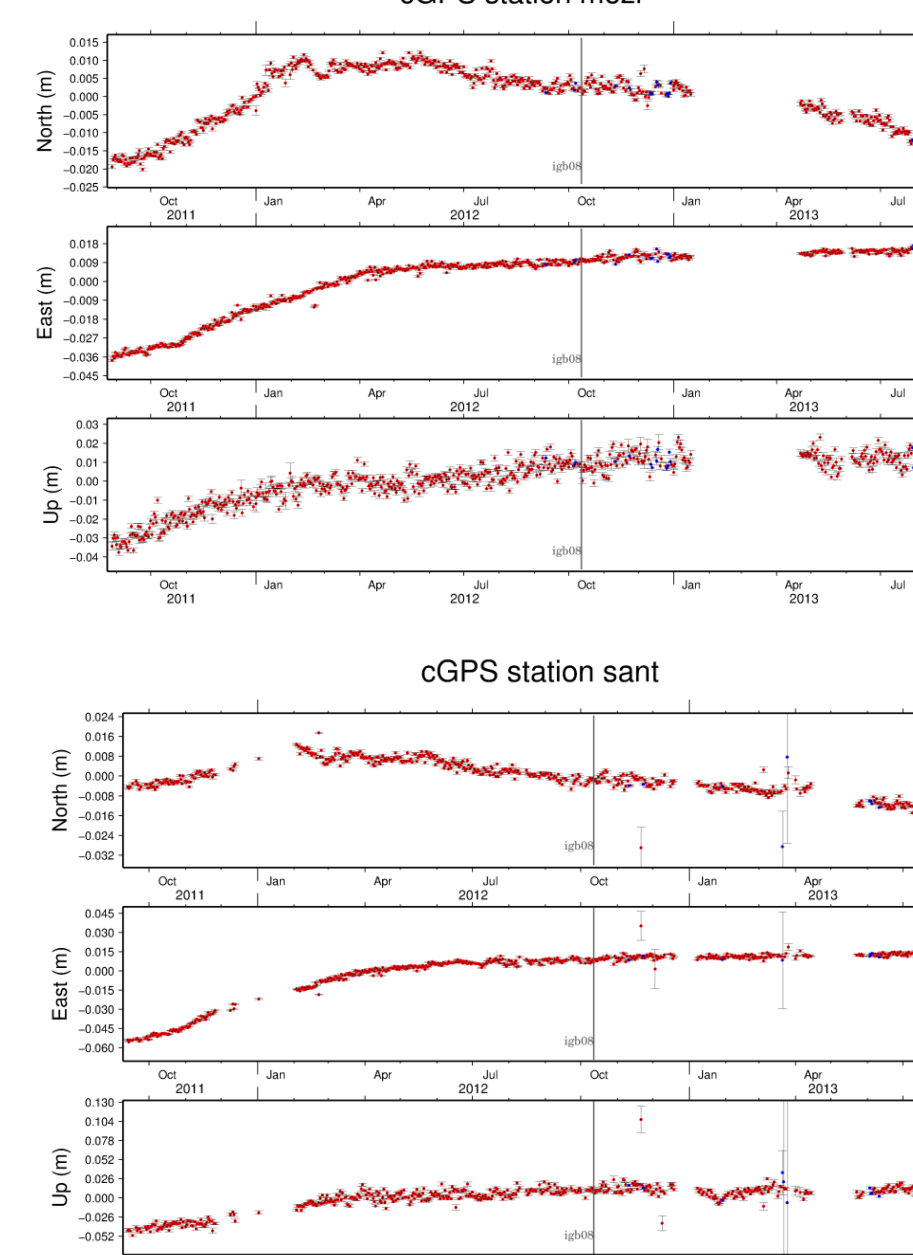
The end of the episode: February/March 2012

DInSAR



Unwrapped differential interferograms in Nea Kameni. While the magnitude of uplift clearly increases for the first three interferograms, in March 2012 the deformation is similar to the one observed in January 2012

GPS



Raw GPS time series up until August 2013 for selected GPS stations

The change in motion around February/March 2012 can be seen

Conclusions

- ✓ Maximum inflation of 150 mm/yr LOS
- ✓ Mogi model seems to be suitable with a source depth of 3.3–6.3 km (in agreement with Newman et al., Parks et al.)
- ✓ Unless a very deep hydrothermal fluid reservoir exists beneath the caldera, this episode was likely to be one of magmatic inflation of the shallow chamber
- ✓ Inflation has diminished since the end of February 2012
- ✓ New phase of relative stability, reduced probability for an imminent volcanic eruption

References

- Papoutsis et al., (2013), Mapping inflation at Santorini volcano, Greece, using GPS and InSAR, *Geophys. Res. Letters*, 40(2): 267-272
- Newman et al., (2012), Recent geodetic unrest at Santorini caldera, Greece, *Geophys. Res. Letters*, 39(6): L06309
- Parks et al. (2012), Evolution of Santorini Volcano dominated by episodic and rapid fluxes of melt from depth, *Nature Geosci.*, 5(10): 749-754
- Hooper et al. (2007), Persistent Scatterer InSAR for Crustal Deformation Analysis, with Application to Volcán Alcedo, Galápagos, *J. Geophys. Res.*, 112, B07407

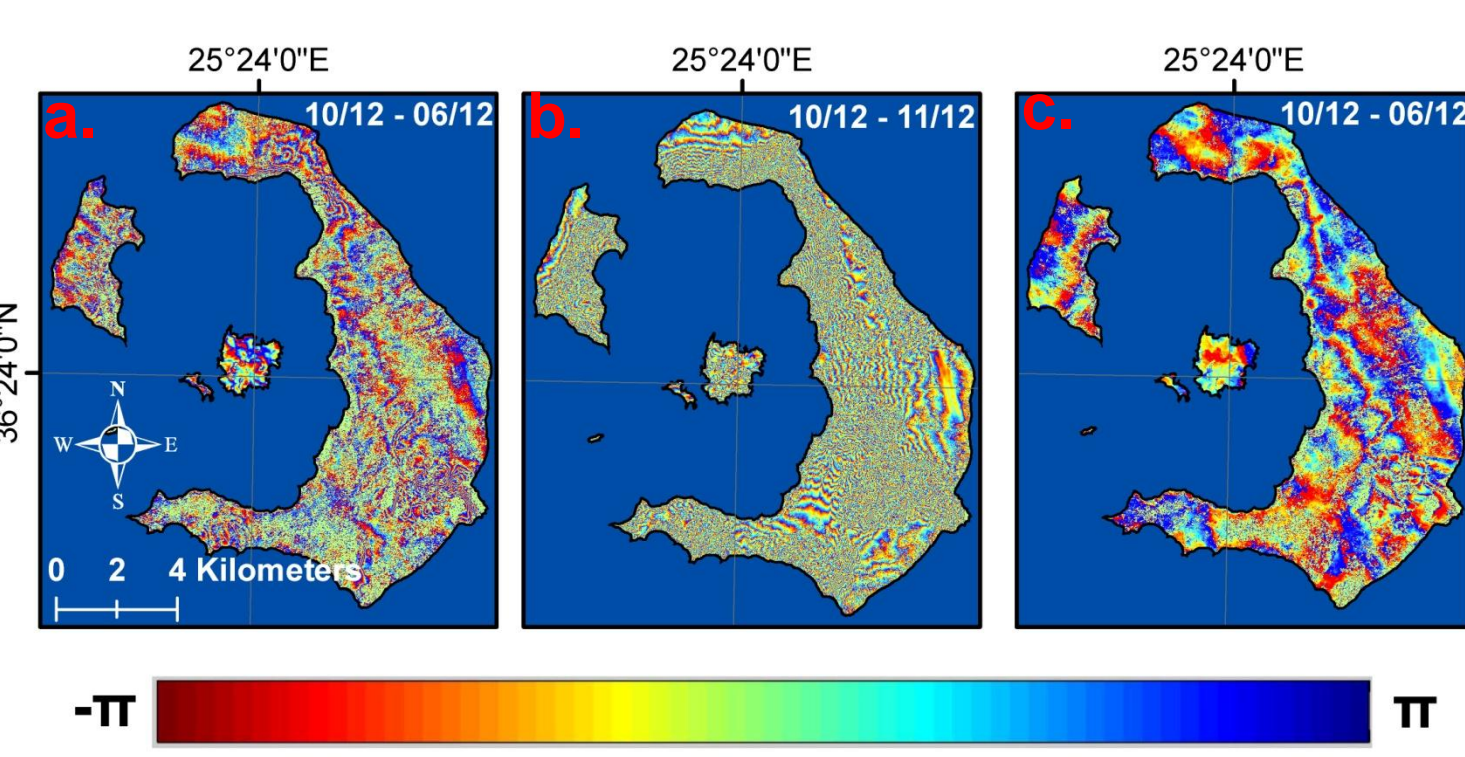
Acknowledgements

- ✓ We acknowledge the ESA provision of Envisat data in the frame of ESA-GREECE AO project 1489OD/11-2003/72.
- ✓ GPS receivers installed and maintained by several bodies: UNAVCO, NTUA, Georgia Tech/University of Patras, COMET/University of Oxford, NOANET/NKUA

Continuous monitoring of Santorini

COSMO-SkyMed

- a. Short perpendicular & long temporal baseline interferogram => deformation
- b. Long perpendicular & short temporal baseline interferogram => DEM
- c. Differential interferogram from 3-pass interferometry



GPS online

Daily GPS solutions for Santorini:
<http://dionysos.survey.ntua.gr/>

BEYOND center of excellence



- ✓ Centre of Excellence for Earth Observation based monitoring of Natural Disasters in south-east Europe
- ✓ <http://www.beyond-eocenter.eu/>
- ✓ June 2013 – 2016, €2.3M EU contribution
- ✓ Beneficiary is the National Observatory of Athens